



## 2 Scaling up ecohydrological processes: Role of surface water flow 3 in water-limited landscapes

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6 [1] In this study, we present a stochastic landscape modeling approach that has the power  
7 to transfer and integrate existing information on vegetation dynamics and hydrological  
8 processes from the small scale to the landscape scale. To include microscale processes like  
9 ecohydrological feedback mechanisms and spatial exchange like surface water flow, we  
10 derive transition probabilities from a fine-scale simulation model. We applied two versions  
11 of the landscape model, one that includes and one that disregards spatial exchange of  
12 water to the situation of a sustainably used research farm and communally used and  
13 degraded rangeland in semiarid Namibia. Our simulation experiments show that including  
14 spatial exchange of overland flow among vegetation patches into our model is a  
15 precondition to reproduce vegetation dynamics, composition, and productivity, as well as  
16 hydrological processes at the landscape scale. In the model version that includes spatial  
17 exchange of water, biomass production at light grazing intensities increases 2.24-fold  
18 compared to the model without overland flow. In contrast, overgrazing destabilizes  
19 positive feedbacks through vegetation and hydrology and decreases the number of  
20 hydrological sinks in the model with overland flow. The buffer capacity of these  
21 hydrological sinks disappears and runoff increases. Here, both models predicted runoff  
22 losses from the system and artificial droughts occurring even in years with good  
23 precipitation. Overall, our study reveals that a thorough understanding of overland flow is  
24 an important precondition for improving the management of semiarid and arid rangelands  
25 with distinct topography.

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### 29 1. Introduction

30 [2] Decisions for the conservation of biodiversity and  
31 sustainable management of natural resources are made for  
32 long time periods and at broad spatial scales [Peters *et al.*,  
33 1997; Miller *et al.*, 2004]. In contrast, our understanding of  
34 the underlying ecological processes (e.g., local water avail-  
35 ability triggering germination rates and plant growth) is  
36 high at fine spatial and temporal scales because most  
37 empirical data are collected for small areas and over a short  
38 duration only [Levin, 1992; Rastetter *et al.*, 2003]. There-  
39 fore, the knowledge from short-term and fine-scale studies  
40 needs to be projected to regional and global scales that are  
41 relevant for decision making [Wessman, 1992].

42 [3] However, extrapolation of information across scales  
43 provides difficulties as we do not know to which extent  
44 spatial exchange, like the movement of surface water by  
45 run-off in water limited environments, affects ecosystem

dynamics at large scales [Levin, 1992; Tongway and  
Ludwig, 1997; Wootton, 2001; Strayer *et al.*, 2003; Urban, 46  
2005]. Omission of these processes may directly affect the 47  
accuracy of predictions [Heuvelink, 1998, Weaver and 48  
Perera, 2004]. Run-off occurs at multiple spatial scales 49  
if rainfall intensity exceeds soil infiltration capacity 50  
[Rango *et al.*, 2006]. Local differences in infiltration 51  
capacity are induced by topography, soil texture and 52  
positive feedback mechanisms between water and vegeta- 53  
tion [Wilcox *et al.*, 2003]. 54  
55

[4] Many arid landscapes are source sink systems, where 56  
plant productivity is determined by surface run-off from 57  
bare areas to vegetated patches. Therefore, theoretical [Noy- 58  
Meir, 1973; Scheffer *et al.*, 2001; Ludwig *et al.*, 2005; 59  
Urban, 2005; Peters and Havstad, 2006] and model inves- 60  
tigations [van de Koppel *et al.*, 2002; van de Koppel and 61  
Rietkerk, 2004] suggest that spatial redistribution of rainfall 62  
by run-off increases the productivity and resilience of arid 63  
ecosystems. 64

[5] Disturbances like unsustainable grazing can disrupt 65  
this fundamental process by changing vegetation structure 66  
and composition [e.g., Ludwig *et al.*, 2005]. The system 67  
may lose its buffer capacity and become less efficient at 68  
trapping run-off, leading to a loss of water. Today, some 69  
20–30% of global drylands [Winand Staring Centre, 1991; 70  
Reid, 2005; Zika and Erb, 2009] and 30% of drylands in 71

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72 Sub-Sahara Africa [Zika and Erb, 2009] are already  
 73 degraded. Maintaining so-called ‘resource conserving’ dry  
 74 lands [Wilcox et al., 2003] will have profound implications  
 75 for management of semiarid and arid rangelands especially  
 76 in the future if climate change will likely lead to a general  
 77 decrease in the grass resource, increase dryland vulnerabil-  
 78 ity to degradation and add an additional pressure on these  
 79 systems already prone to degradation.

80 [6] In the past, spatial transition based models like  
 81 Markov chains have often been used to explore vegetation  
 82 dynamics over long time periods and on large scales [e.g.,  
 83 Baker, 1989; Turner, 1989; Acevedo et al., 1995; Balzter,  
 84 2000; Logofet and Lesnaya, 2000; Urban, 2005]. They are  
 85 based on stochastic processes and can be parameterized by  
 86 estimating transition probabilities between discrete states of  
 87 the observed system. Most previous studies utilized data  
 88 sampled from field surveys, existing maps, aerial photo-  
 89 graphs or satellite images to estimate transition probabilities  
 90 [e.g., Muller and Middleton, 1994; Brown et al., 2000;  
 91 Jenerette and Wu, 2001; Weng, 2002], and only very few  
 92 studies exist that make use of a fine-scale model to drive a  
 93 landscape model [e.g., Acevedo et al., 1995]. Generally,  
 94 uncertainty in these studies remains relatively high because  
 95 spatial exchange data such as surface water run-off are not  
 96 considered. Furthermore, as very few studies do exist that  
 97 describe processes at the landscape scale mechanistically  
 98 [e.g., Acevedo et al., 1995], data is limited and transition  
 99 probabilities are often derived from short-term data [Baker,  
 100 1989; Urban, 2005].

101 [7] In this study we developed a method to transfer and  
 102 integrate existing information on vegetation dynamics and  
 103 hydrological processes in arid rangelands between spatial  
 104 scales. We used a small-scale simulation model [Popp et al.,  
 105 2009], that was developed to investigate the relative impact  
 106 of small-scale soil-plant interactions on vegetation dynam-  
 107 ics at the hillslope scale, to derive transition probabilities  
 108 between different vegetation states, productivity coefficients  
 109 of different vegetation types, and hydrological parameters  
 110 for a model operating at the landscape scale. We applied this  
 111 modeling approach on a dwarf shrub savannah with distinct  
 112 topography in arid southern Namibia to assess the role of  
 113 surface water flow at the landscape scale for low and high  
 114 grazing intensities.

## 115 2. Material and Methods

116 [8] Two variants of a stochastic and spatially explicit  
 117 landscape model were implemented on the basis of Mar-  
 118 kovian modeling that simulates annual biomass production  
 119 of a dwarf shrub savannah with distinct topography in arid  
 120 southern Namibia (Karas Region). One version simulates  
 121 lateral exchange of surface water, whereas explicit consid-  
 122 eration of overland flow is eliminated in the second version.  
 123 We used a small-scale simulation model (Topographical  
 124 Management (TOPMAN)) [Popp et al., 2009], that was  
 125 developed to investigate the relative impact of small-scale  
 126 soil-plant interactions on vegetation dynamics at the hill-  
 127 slope scale, to derive data which is handed over to the  
 128 landscape model. This mechanistic approach guarantees that  
 129 data collected and processes estimated at smaller scales are  
 130 included in our application. Elevation of the landscape’s  
 131 grid cells was parameterized by remotely sensed digital

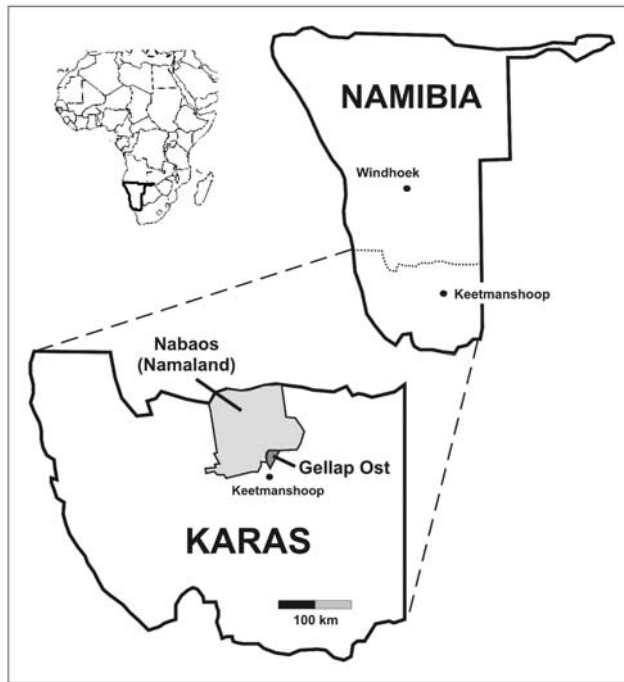
elevation models (DEM). We computed annual productivity 132  
 of the four most abundant vegetation types for two con- 133  
 trasting land management systems: a sustainably used 134  
 research farm (for Karakul sheep breeding) and communal 135  
 farming land, on which livestock grazing pressure is not 136  
 controlled. 137

### 2.1. Study Area 138

[9] The study area is located in the Nama Karoo, southern 139  
 Namibia (Figure 1). Vegetation cover and productivity are 140  
 low and depend on erratic and highly variable rainfall 141  
 (annual mean: 150 mm). The main topographical features 142  
 of the study area are flat regions, as well as regions with 143  
 gentle and precipitous slopes [Kuiper and Meadows, 2002]. 144  
 Perennial grasses (e.g., *Stipagrostis uniplumis*) dominate the 145  
 herbaceous vegetation if the rangeland is in good condition 146  
 but are replaced by annual grasses (such as *Schmidtia* 147  
*kalahariensis*) and unpalatable shrubs (like *Rhigozum tri-* 148  
*chotomum*) when rangeland is heavily utilized [Kuiper and 149  
*Meadows, 2002*]. The most important land use in the 150  
 communal area (Nabaos) and the Research Station (Gellap 151  
 Ost) is small stock farming. Gellap Ost has 160 purposely 152  
 understocked camps (0.05 SSU – small stock unit ha<sup>-1</sup>), 153  
 where animals graze in a rotational system. Resting periods 154  
 of camps (no grazing) of at least one year prevent over- 155  
 grazing [Kuiper and Meadows, 2002]. In contrast, Nabaos is 156  
 managed under a communal land tenure system where 157  
 livestock movement in the area is not controlled and over- 158  
 stocking (0.2 SSU ha<sup>-1</sup>) has a strong impact on the 159  
 rangeland resource. 160

### 2.2. Landscape Model 161

[10] The grid based landscape model simulates vegetation 162  
 dynamics and interlinked hydrological processes of a dwarf 163  
 shrub savannah with distinct topography in arid southern 164  
 Namibia. Data on these dynamics and processes are derived 165  
 from a small-scale (spatial resolution: 3 × 3m cells, 33 × 33 166  
 cells ≈ 1 ha) and spatially explicit simulation model (TOP- 167  
 MAN) [Popp et al., 2009]. The area simulated by the small- 168  
 scale model (1 ha) was used as the spatial resolution (cell 169  
 size) for the landscape model. Each cell is specified by its 170  
 position within the grid, elevation, productivity, vegetation 171  
 composition and water availability (composed of precipita- 172  
 tion and surface run-off). Generally, water availability is 173  
 composed of the interplay of precipitation, overland flow, 174  
 deep percolation and total evapotranspiration, which is 175  
 composed of vegetation interception, evaporation from the 176  
 soil and transpiration by plants [Wilcox et al., 2003]. 177  
 However, the landscape model does not consider all of 178  
 these processes explicitly: Recent research suggests that 179  
 deep drainage and groundwater recharge do not occur in 180  
 many arid and semiarid landscapes. This is mainly because 181  
 the potential for increased deep percolation during wet years 182  
 is countered by an increase in the density of plants, a 183  
 concomitant increase in the density of deep roots, and 184  
 possibly an increase in the depth of the root zone as well 185  
 [e.g., Walvoord et al., 2002; Seyfried et al., 2005]. Second, 186  
 we disregarded vegetation interception because, in contrast 187  
 to humid landscapes [e.g., Crockford and Richardson, 188  
 2000], vegetation cover in arid and semiarid regions has 189  
 comparatively little effect on vegetation interception in arid 190  
 environments [Huxman et al., 2005]. Third, we did not 191



**Figure 1.** Location of the two case studies commercial research farm Gellap Ost and communal rangeland Nabaos.

192 include small-scale soil-plant interactions that affect  
193 infiltration and evaporation from the soil because these  
194 processes are already brought in via the small-scale model  
195 [see Popp *et al.*, 2009].

196 [11] Dynamics of the simulated vegetation types (peren-  
197 nial grass, annual forbs, dwarf shrubs and shrubs) on the  
198 cell level are based on the concept of state and transition  
199 models [Westoby *et al.*, 1989]. These models provide a  
200 relatively simple, management-oriented way to classify land  
201 condition (state) and to analyze the impact of factors  
202 that might cause a shift to another state (transition). The  
203 stochastic process of state and transition models and  
204 resulting forecasting of land cover change can be projected  
205 by Markov chain models [Markov, 1907]. To construct a  
206 Markov chain, we first identified vegetation states for the  
207 research area jointly with Namibian rangeland experts  
208 (including farmers and extension officers). The definition  
209 of these six states is related to the percentage cover of  
210 shrubs and perennial grasses (Table 1).

211 [12] Subsequently, we calculated annual transition  
212 probabilities between these states from simulations using  
213 TOPMAN [Popp *et al.*, 2009] for different classes of slope,  
214 rainfall and land use. For slope, we defined four classes: flat  
215 (<6%), gentle (6–10%), steep (11–15%) and precipitous  
216 (>15%). Water availability was classified in four categories,  
217 relative to the long-term mean of 150 mm: 1 is poor  
218 (<100 mm), 2 is moderate (100–139 mm), 3 is good  
219 (140–179 mm), and 4 is very good (>180 mm). Finally,  
220 we defined three categories of land use: no grazing, under-  
221 stocked grazing (0.05 small stock unit (SSU) ha<sup>-1</sup>) and  
222 overstocked (0.2 SSU ha<sup>-1</sup>) grazing. In contrast to the  
223 communal and ‘overstocked’ land tenure system at Nabaos  
224 where livestock movement is not controlled, the grid cells at  
225 the purposely understocked and rotationally grazed research  
226 farm (Gellap Ost) are alternately treated by ‘no grazing’ and

‘understocked grazing’. Within each time step (1 year), the  
227 landscape model calculates the following modules in the  
228 given order: water availability, vegetation dynamics and  
229 productivity (Figure 2). Each model is explained below. 230

### 2.3. Model Initialization 231

[13] Calculation of the slope ( $s$ ) of each 1 ha cell is based  
232 on elevation ( $e_c$ ) and side length ( $l = 100$  m) of the  
233 respective cell and elevation of the neighboring cells ( $e_{nc}$ ) 234

$$s = \frac{e_c - e_{nc}}{l}. \quad (1)$$

Elevation was initialized with a digital elevation model 236  
(DEM), derived by remote sensing based radar data. The  
237 raster DEM is processed interferometrically from SRTM C  
238 Band data [Jensen, 2000] with an original spatial resolution  
239 of 88 m  $\times$  88 m in  $x$  and  $y$  direction and 1 m resolution of  
240 the altitude ( $z$  direction). The data have been preprocessed,  
241 applying a 3  $\times$  3 kernel low-pass filter to reduce radar  
242 system inherent errors, caused by signal noise (“salt and  
243 pepper effect”), and shadow effects [Lewis, 1976]. 244  
Application of this filter leads to a smoothing of high  
245 contrast image areas. To fit the 100  $\times$  100 m cell size of the  
246 landscape model, the DEM subset of the study area has been  
247 resampled, using a Nearest Neighbor algorithm. 248

[14] For initialization of the vegetation, we used the  
249 vegetation structure of an undisturbed dwarf shrub savan-  
250 nah. Since little is known about this vegetation structure we  
251 assumed an undisturbed coexistence of perennial grasses  
252 and woody vegetation. Thus initial vegetation condition for  
253 each cell was set to state 3 (compare Table 1). 254

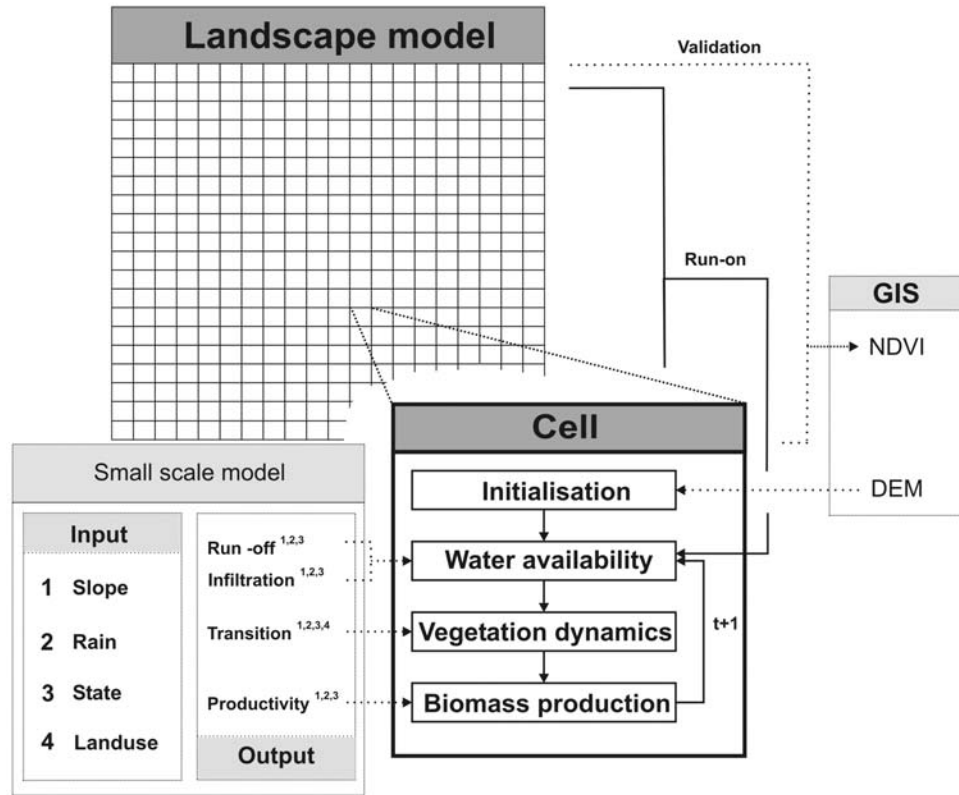
### 2.4. Vegetation Dynamics 255

[15] We used a state and transition approach to simulate  
256 the vegetation dynamics. Transition probabilities between  
257 the vegetation states were calculated as Markovian sto-  
258 chastic processes [Markov, 1907]: a state at time  $t$   
259 depends on the state at time  $t - 1$  and the impact of  
260 the exogenous factors water availability, slope and present  
261 land use. A  $m \times m$  transition matrix ( $P$ ) contains the  
262 conditional probabilities  $p_{ij}$  that a cell in state  $i$  at time  $t$   
263 will transition to state  $j$  at time  $t + 1$ .  $P$  is row  
264 standardized, such that the sum of transition probabilities  
265 from a given state is always equal to one. We derived  
266 transition probabilities by calculating transition probabili-  
267 ties of the process-based small-scale simulation model. To  
268 gain these values we ran the small-scale model for 100  
269 years with 500 repetitions. 270

**Table 1.** Definition of Vegetation States for the Research Area<sup>a</sup> t1.1

	Cover <sub>PG</sub> (%)	Cover <sub>W</sub> (%)	
State 1	30 – 100	40 – 100	t1.3
State 2	0 – 29	40 – 100	t1.4
State 3	50 – 100	0 – 40	t1.5
State 4	10 – 49	0 – 40	t1.6
State 5	0 – 9	10 – 40	t1.7
State 6	0 – 9	0 – 9	t1.8

<sup>a</sup>Columns refer to cover of the respective vegetation type (PG is  
perennial grass and W is woody vegetation). Rows refer to vegetation state,  
enumerated by 1–6. t1.9



**Figure 2.** Visualization of mechanistic upscaling approach and simplified flowchart. We used a small-scale simulation model to derive data for vegetation dynamics, productivity, and linked hydrological processes of the landscape model. Elevation of the landscape's grid cells is initialized by remotely sensed digital elevation models (DEM). For validation, simulated annual biomass production is compared with remotely sensed estimates of annual biomass production (normalized differential vegetation index (NDVI)). Solid lines represent processes within the landscape model (flowchart). Dotted lines illustrate data flow between different disciplines and scales. Numbers refer to basic attributes of the small-scale model (input) affecting transferred data for the landscape model (output).

271 [16] Generally, we assume that (1) transition probabilities  
 272 are constant over time and (2) transitions are spatially  
 273 independent. An approach to model nonstationarity (i.e.,  
 274 variability in space and time of transition probabilities) is to  
 275 switch between different stationary matrices [Rejmanek et  
 276 al., 1987]. To consider variability in space and time of  
 277 transition probabilities related to water availability, slope,  
 278 and land use option, we generated transition matrices for all  
 279 combinations of these exogenous factors.

## 280 2.5. Productivity

281 [17] Annual phytomass production of the four vegetation  
 282 types is dependent on the cells slope, present vegetation  
 283 state, and water availability. For each vegetation type we  
 284 calculated productivity coefficients from multiple linear  
 285 regressions of biomass on slope, vegetation state and water  
 286 availability simulated by the small scale model.

## 287 2.6. Water Availability

288 [18] We implemented two versions of the landscape  
 289 model. One version disregards lateral exchange of surface  
 290 water, while the other explicitly considers overland flow. In  
 291 the first version, water availability ( $W_C$ ) is dependent on  
 292 annual precipitation ( $p_C$ ). In the case of overland flow

(second version),  $W_C$  is not only related to annual precip- 293  
 itation ( $p$ ) but also to run-on ( $r_C$ ) from neighboring cells and 294  
 can be expressed by 295

$$W_c = p + r_c, \quad (2)$$

where  $p$  is homogeneous for all cells, whereas  $r_C$  is based 297  
 on a cell specific capacity to absorb run-on ( $ir_C$ ) and the 298  
 contribution by run-off from neighboring cells ( $r_{NC}$ ) 299

$$r_C = r_{NC} - ir_C. \quad (3)$$

We used an iterative algorithm to calculate surface water 301  
 flow for each simulated year: in the first step, each cell's 302  
 $ir_{C,0}$  and  $r_{NC,0}$ , calculated by the small-scale model, are 303  
 based on rain class as well as a cell's slope class and current 304  
 vegetation state. 305

[19] In each following iterative step  $r_{NC,i}$  is updated until 306  
 infiltration of a cell is saturated ( $ir_{C,i} = 0$ ) and until no more 307  
 cells pass flow ( $r_{NC,i} = 0$ ). For the next iterative step  $ir_{C,i+1}$  308  
 is actualized by 309

$$ir_{C,i+1} = ir_{C,i} - r_{C,i}. \quad (4)$$

311 We used multiple flow direction methods to estimate surface  
312 water flow directions across cells [Quinn *et al.*, 1991;  
313 Tarboton, 1997]. Thus, cells with run-off allocate water  
314 fractionally to each lower neighbor cell in proportion to the  
315 respective slope.

## 316 2.7. Small-Scale Simulation Model

317 [20] Transition probabilities and phytomass production  
318 are derived from a small-scale simulation model [Popp  
319 *et al.*, 2009]. The spatially explicit and individual based  
320 model simulates the vegetation dynamics of a 1 ha area  
321 ( $100 \times 100$  m,  $33 \times 33$  cells). Cell size is  $3 \text{ m} \times 3 \text{ m}$ , which  
322 corresponds to the maximum observed diameter of a shrub.  
323 Herbaceous vegetation (perennial grass and annuals) are  
324 treated as matrix plants, and a cell is either occupied or not.  
325 Woody plants are simulated individually and each cell  
326 contains a list of woody plant individuals. Slope angle is  
327 included by decreasing elevation values of the cells toward  
328 one side of the landscape. For each cell, water availability,  
329 establishment, aboveground biomass production, grazing  
330 and mortality are simulated in annual time steps.

### 331 2.7.1. Water Availability

332 [21] Cell-specific water availability in the various  
333 vegetation types is influenced by rainfall, run-off, run-on,  
334 evapotranspiration and competition. Surface run-off in  
335 semiarid and arid regions occurs primarily as infiltration  
336 excess overland flow from higher to lower areas controlled  
337 by infiltration characteristics of the soil surface rather than  
338 the storage capacity of the soil [Wilcox *et al.*, 2003].  
339 Therefore, infiltration rates are related to the cell's soil  
340 texture [Bergkamp, 1998], vegetation cover [Puigdefabregas,  
341 2005] and slope [Chaplot and Le Bissonnais, 2000]. Surface  
342 run-off occurring at cells with lowest elevation values at  
343 the edge of the simulated landscape leaves the system.  
344 Moreover, soil texture and vegetation cover have an impact  
345 on evaporation [LeHouerou, 1984; Snyman, 2000], reducing  
346 soil water content in the upper soil layer. Competitive effects  
347 of vegetation (i.e., vegetation-specific transpiration) reduce  
348 the water availability for establishment and, in cells with  
349 overlapping root systems, vegetation has a competitive effect  
350 on the neighboring vegetation types [Callaway and Walker,  
351 1997]. To simplify the hydrological processes in our  
352 modeling approach we assume that all infiltrated water is  
353 evaporated or transpired.

### 354 2.7.2. Establishment

355 [22] In arid and semiarid environments, sufficient  
356 moisture [O'Connor, 1994] and the availability of seeds  
357 [O'Connor and Pickett, 1992] are the main conditions for  
358 successful plant establishment. Furthermore, limited food  
359 for livestock increases the probability that livestock will  
360 feed on seedlings [Carrick, 2003]. Therefore, the cells'  
361 probability of successful establishment of perennial vegeta-  
362 tion (woody plants and perennial grasses) is determined by  
363 site-specific probabilities of seed and water availability, as  
364 well as the probability to survive grazing. Annuals, produc-  
365 ing large numbers of seeds and persistent seed banks  
366 [Veenendaal *et al.*, 1996] are only restricted by water  
367 availability and grazing pressure.

### 368 2.7.3. Growth

369 [23] Biomass production of the herbaceous vegetation is  
370 related to annual water availability for the two matrix plant  
371 types: annual and perennial grasses. For both vegetation

types, we use a growth coefficient derived from rainfall- 372  
grass production relationships of various southern African 373  
savannah regions [Higgins *et al.*, 2000]. In contrast, annual 374  
biomass production of woody plant individuals depends on 375  
the impacts of water availability and is related to current 376  
height performance. 377

### 2.7.4. Grazing 378

[24] The model simulates grazing and browsing on 379  
herbaceous and woody vegetation. Vegetation types differ 380  
in their palatability for grazers and browsers. What and to 381  
which amount a plant's biomass will be consumed depends 382  
on its specific palatability as well as the relation of available 383  
biomass in the landscape and that necessary for forage. 384

### 2.7.5. Mortality 385

[25] Survival of perennial plants is environmentally 386  
determined by the availability of water and the impact of 387  
grazing by livestock [Milton and Dean, 2000]. We related 388  
the survival probability to these factors (water and grazing), 389  
since disturbances such as drought or overgrazing strongly 390  
influence productivity. Within the group of perennial plants, 391  
differences among species in disturbance tolerance are 392  
associated with physiological adaptations to disturbance. 393

## 2.8. Simulation Analysis 394

[26] We used the landscape model to simulate vegetation 395  
dynamics and productivity of perennial grass, annuals, 396  
dwarf shrubs and shrubs for 150 years. Effects of connec- 397  
tivity and spatial explicitness of overland flow on dynamics 398  
and productivity of the most abundant vegetation types 399  
was assessed for a purposely understocked and rotationally 400  
grazed research farm (Gellap Ost) ( $0.05 \text{ SSU ha}^{-1}$ ) and 401  
a communal and overstocked range land (Nabaos) 402  
( $0.2 \text{ SSU ha}^{-1}$ ). For model analysis, we used the years 403  
1985 to 2000, as remotely sensed data is available only for 404  
this time period, and initialization effects could be excluded. 405  
Due to the stochastic processes in the model, no single run 406  
is representative. Therefore, we initiated 50 repeats for each 407  
model type and scenario. 408

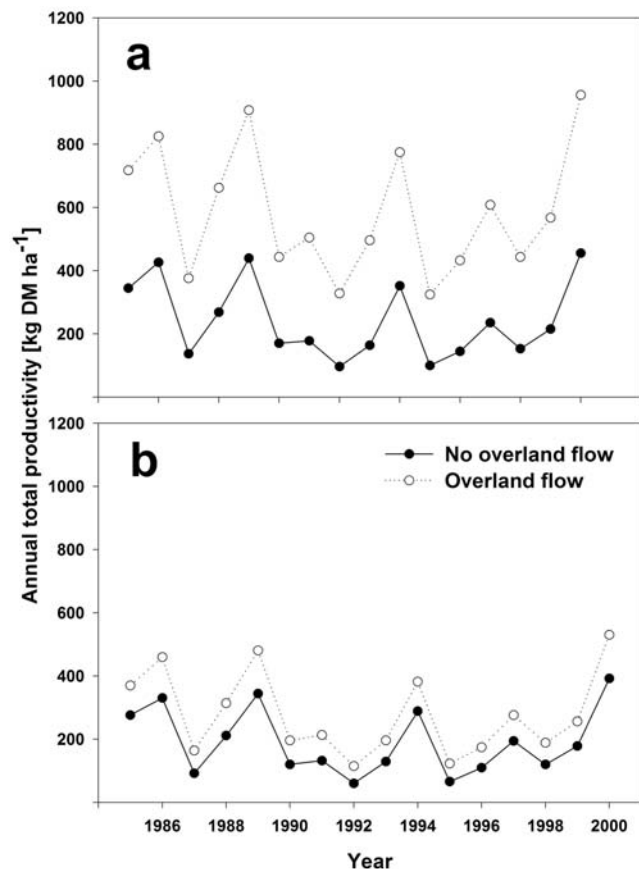
## 2.9. Model Validation 409

[27] We compared simulation results of annual phytomass 410  
production for both study sites (low grazing impact versus 411  
high grazing impact) with remotely sensed Normalized 412  
Difference Vegetation Index (NDVI) as an indicator for 413  
interannual variability. Because NDVI indicates relative 414  
values we also compared our model results with field data 415  
on biomass production and cover for perennial grasses. 416

[28] Remotely sensed NDVI is strongly correlated with 417  
phytomass production [e.g., Tucker *et al.*, 1986; Prince and 418  
Goward, 1996; Yang and Prince, 2000; Wessels *et al.*, 419  
2006]. NDVI is calculated from the red and near-infrared 420  
channels from multispectral remote sensing imagery [Tucker 421  
and Choudhury, 1987] 422

$$NDVI = \frac{NIR - R}{NIR + R}, \quad (5)$$

where NIR is the reflectance in the near infrared band 424  
( $0,72 - 1,10 \mu\text{m}$ ) and R the reflectance in the red band 425  
( $0,58 - 0,68 \mu\text{m}$ ). NDVI time series data for 1985 to 2001 426  
were obtained from NOAA/NASA Pathfinder Land data 427  
archive (PAL). However, due to the failure of NOAA-11, no 428



**Figure 3.** Time series of annual total productivity. Medians of 500 simulation replicates are shown. Annual total productivity at (a) low grazing intensity (Gellap) is strongly affected by overland flow. Low impact of overland flow can be found at the scenario with (b) high grazing intensity (Nabaos). Black circles represent scenarios without overland flow, and white circles represent scenarios with overland flow.

429 NDVI data were available between July and December  
430 1994 (<http://www2.ncdc.noaa.gov/docs/gviug/html/c2/sec2-0.htm>).

432 [29] We have used the seasonal integral or accumulated  
433 NDVI (I-NDVI) that was calculated for each sampling  
434 domain from seasonal summations (October to September)  
435 of differences between NDVI and minimum NDVI from the  
436 seasons 1985–1986 to 2000–2001 [Holm *et al.*, 2003].  
437 Reference values of NDVI have been calculated for both  
438 study sites over an area of 8 km<sup>2</sup> each.

### 440 3. Results

#### 441 3.1. Annual Total Phytomass

442 [30] Annual total phytomass simulated by the model  
443 without overland flow ranged from 475 t per ha at the  
444 scenario with low grazing intensity (Gellap Ost, Figure 3a)  
445 in 2000 to 79 t per ha at the scenario with high grazing  
446 intensity (Nabaos, Figure 3b) in 1992. In contrast, in all  
447 simulations using the model including overland flow,  
448 annual total phytomass ranged from 983 t per ha at low

grazing intensity in 2000 (Figure 3a) to 112 t per ha at high  
grazing intensity in 1992 (Figure 3b).

[31] Averaged over a period of 15 years in the model  
without overland flow, total phytomass production simulat-  
ed by the scenario with low grazing intensity exceeds total  
phytomass production simulated by the scenario with high  
grazing intensity 1.32-fold. In the model including overland  
flow, the 15 year average of total phytomass production for  
low grazing intensity exceeds the total phytomass produc-  
tion for the high grazing intensity scenario 2.24-fold.

[32] Linear regression analysis for the model without  
overland flow indicated that total phytomass at the scenario  
with low grazing intensity ( $R^2 = 0.97$ ,  $p < 0.001$ ) and the  
scenario with high grazing intensity ( $R^2 = 0.99$ ,  $p < 0.001$ )  
increased with annual precipitation. In the model that  
includes overland flow, total phytomass equally increased  
with annual precipitation for the scenario with low grazing  
intensity and for high grazing intensity (both  $R^2 = 0.98$ ,  $p <$   
 $0.001$ ). In the model that includes overland flow, total  
phytomass production is 2.61-fold higher at the scenario  
with low grazing intensity and 1.54-fold higher at the  
scenario with high grazing intensity, than that of the model  
without overland flow.

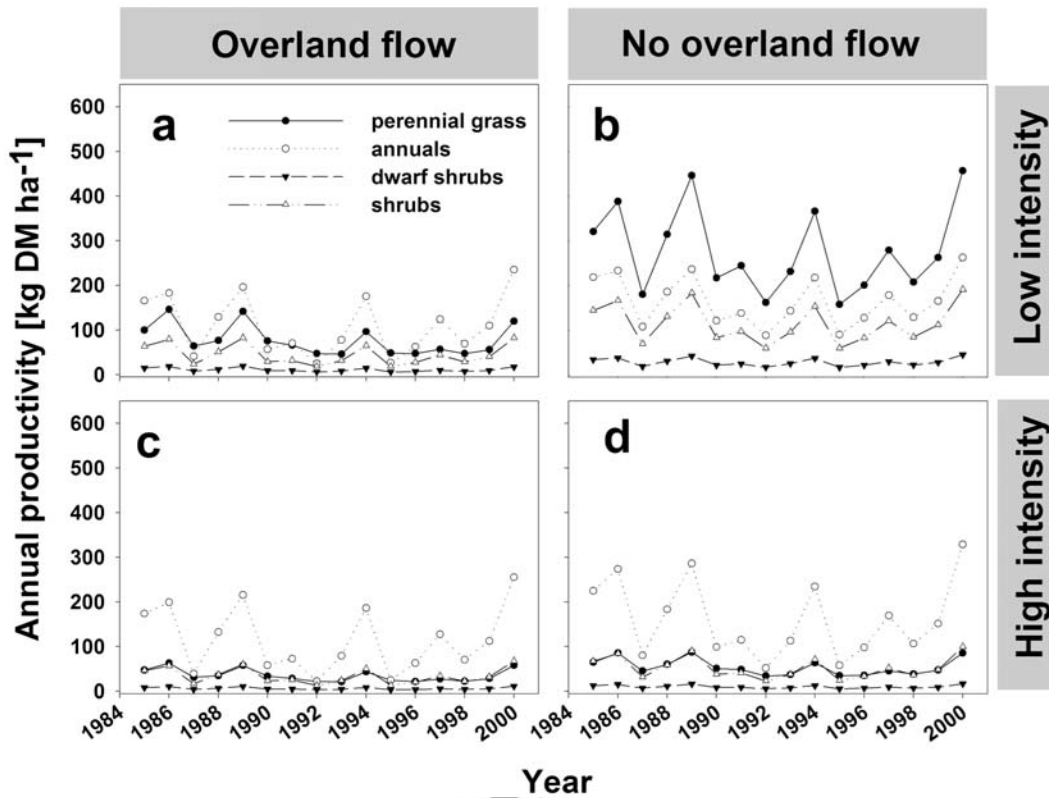
#### 472 3.2. Productivity of Vegetation Types

[33] In the model without overland flow, productivity of  
annuals contributed 45% to the mean total productivity,  
followed by perennial grass (31%), shrubs (18%) and dwarf  
shrubs (6%) for the scenario with low grazing intensity  
(Figure 4a). At the scenario with high grazing intensity  
(Figure 4c), the proportion in mean annual total productivity  
was clearly decreased for perennial grasses (19%) and dwarf  
shrubs (3%), and dominance was shifted toward annuals  
(57%) and shrubs (22%).

[34] In the model that includes overland flow, it is not  
only total annual phytomass production that is affected by  
spatial exchange of surface water, precipitation and land  
use; it is also productivity of the four most abundant  
vegetation types (perennial grass, annuals, dwarf shrubs  
and shrubs). Including overland flow in simulating dynamics  
and productivity of these vegetation types has the strongest  
impact at the scenario with low grazing intensity (Figure 4b):  
perennial grass contributes with 47% most to the mean total  
productivity, followed by annuals (28%), shrubs (19%) and  
dwarf shrubs (6%). At the scenario with high grazing  
intensity (Figure 4d), no effect of overland flow on the  
proportion of vegetation types in mean total productivity  
could be identified. However, the model excluding overland  
flow clearly indicates a decrease in the proportion in mean  
annual total productivity for perennial grasses (19%) and  
dwarf shrubs (3%), and dominance was shifted toward  
annuals (60%) and shrubs (18%).

#### 500 3.3. Disturbance and Overland Flow

[35] Disturbance in the form of overgrazing can have a  
strong impact on lateral exchange of surface water  
(Figure 5). Light grazing intensities lead to 83% of run-on  
cells in the total number of cells (% of area) as well as high  
mean run-on in these cells (85%) for the years 1985 to  
2000. In contrast, the proportion of run-on cells to total area  
decreases to 63% with low mean run-on of 40% at the  
scenario with high grazing intensity.



**Figure 4.** Time series of annual productivity for the simulated vegetation types. Medians of 500 simulation replicates are shown. Different management scenarios ((a and c) Gellap and (b and d) Nabaos) are shown. Figures 4a and 4b show results for simulation models without overland flow, and Figures 4c and 4d refer to simulation models including overland flow. Black circles refer to perennial grass, white circles refer to annual forbs, black triangles refer to dwarf shrubs, and white triangles refer to shrubs.

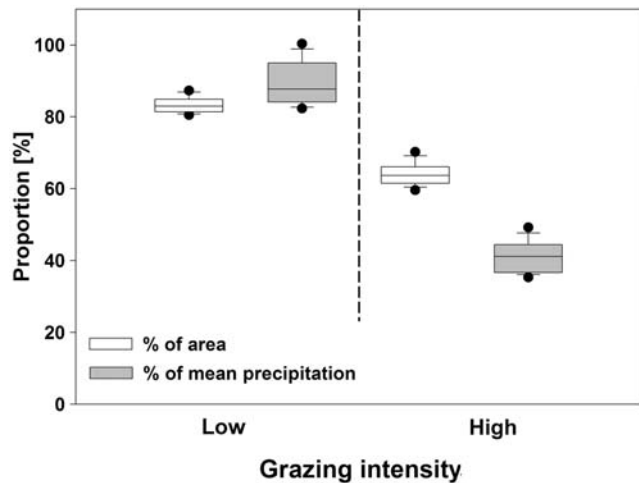
509 **3.4. Validation of Simulated Phytomass Production**

510 **3.4.1. Estimated I-NDVI**

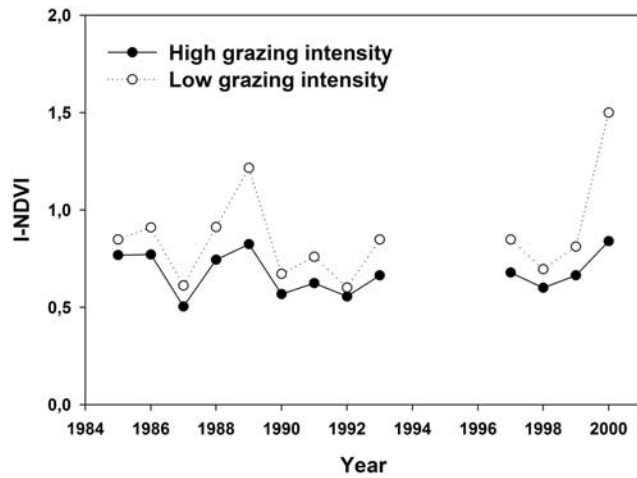
511 [36] Vegetation at the communal rangelands (Nabaos) had  
 512 lower I-NDVI than vegetation at the research farm (Gellap  
 513 Ost) across all growth seasons (Figure 6). The lowest value  
 514 of 0.5 at the communal rangelands in 1987 contrasts with  
 515 the highest value of 1.6 at the research farm in 2000.  
 516 Averaged over the time span of 15 years, I-NDVI measured  
 517 at the research farm exceeds I-NDVI measured at the  
 518 communal rangelands 1.27-fold. Linear regression analysis  
 519 indicated that I-NDVI is correlated with annual precipitation  
 520 at the research farm ( $R^2 = 0.81$ ,  $p < 0.001$ ) and the  
 521 communal rangelands ( $R^2 = 0.91$ ,  $p < 0.001$ ).

522 **3.4.2. Comparison of Simulated Phytomass  
 523 and Estimated I-NDVI**

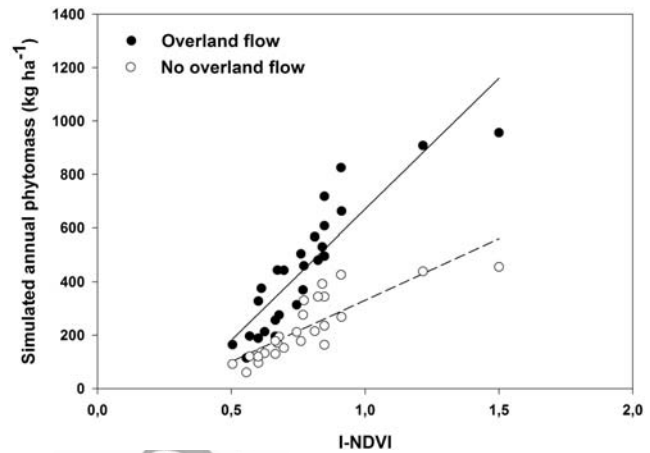
524 [37] We compared simulated total annual phytomass with  
 525 remotely sensed estimates of annual phytomass production  
 526 (I-NDVI) to test if the landscape model displays vegetation  
 527 dynamics in a simplified but realistic way. The linear  
 528 regression relationship between simulated total phytomass  
 529 and I-NDVI (Figure 7) accounted for little more variance at  
 530 the model with overland flow ( $R^2 = 0.79$ ,  $p < 0.001$ ) than  
 531 for the model version without overland flow ( $R^2 = 0.69$ ,  $p <$   
 532  $0.001$ ). However, the slopes of the linear regression rela-  
 533 tionship indicate that total productivity simulated by the  
 534 model with overland flow ( $y = -308 + x * 978$ ) exceeds



**Figure 5.** Importance of overland flow under different grazing intensities. At the scenario with low grazing intensity (left side) overland flow leads to high proportion of run-on cells to total area (white box plots) and high values of mean run-on (gray box plots). Mean run-on is given as a percent in total annual rain. In contrast, proportion of run-on cells as well as mean proportion of run-on at the scenario with high grazing intensity (right side) display low values.



**Figure 6.** Time series of integrated normalized differential vegetation index (I-NDVI). I-NDVI was calculated for each sampling domain from seasonal summations (October–September) of differences between NDVI and minimum NDVI from the seasons 1985–1986 to 2000–2001. Black circles represent the scenario with low grazing intensity (Gellap), and white circles represent the scenario with high grazing intensity (Nabaos).



**Figure 7.** Comparison of simulated phytomass and measured I-NDVI. For both simulation methods and both management methods modeled total phytomass in each year from 1985 to 2000 were compared with remotely sensed indices of phytomass (I-NDVI). I-NDVI images cover 8km<sup>2</sup> of both the research farm with low grazing intensity (Gellap) and communal rangeland with high grazing intensity (Nabaos).

535 productivity simulated by the model without overland flow  
536 ( $y = -129 + x * 460$ ).

537 **3.4.3. Comparison of Simulated Phytomass**  
538 **and Cover With Measured Field Data**

539 [38] Perennial grass cover and biomass production simu-  
540 lated by the model with overland flow for the year 2001  
541 matched the field observations [Prinsloo and Bester, 2003;  
542 Wolkenhauer, 2004] more effectively than simulation results  
543 from the model version without overland flow (Table 2).

545 **4. Discussion**

546 [39] The aim of our study was to present a method to  
547 transfer and integrate existing information on vegetation  
548 dynamics and hydrological processes between spatial  
549 scales. Combining technologies of remote sensing technol-  
550 ogy and stochastic modeling, we are able to successful  
551 extrapolate vegetation dynamics, composition and produc-  
552 tivity to an order of magnitude one hundred times greater  
553 than the original scale. Using a small-scaled simulation  
554 model, the influence of exogenous and endogenous varia-  
555 bles (vegetation state, precipitation, management and  
556 topography) on transition probabilities and phytomass  
557 production were estimated.

558 [40] In the past, transition probabilities of stochastic  
559 landscape models were mainly estimated by using data  
560 from observations, and measured from empirical studies,  
561 aerial photography and satellite images [e.g., Muller and  
562 Middleton, 1994; Brown et al., 2000; Jenerette and Wu,  
563 2001; Weng, 2002]. Uncertainty in these studies remained  
564 relatively high because data was limited, i.e., transition  
565 probabilities were derived from short-term data [Baker,  
566 1989]. The few studies that use a fine-scale model to drive  
567 a landscape model [e.g., Acevedo et al., 1995] did not  
568 include microscale processes like ecohydrological feedback  
569 mechanisms and spatial exchange like surface water flow.

[41] The simulation results of our landscape model show 570  
571 that the explicit consideration of surface water flow can  
572 have a strong impact on vegetation dynamics, composition  
573 and productivity at the landscape scale. At low grazing 574  
575 intensity, a high number of run-on cells with high infiltra-  
576 tion capacities serve as sinks for input by surface water  
577 flow. Spatial exchange of surface water among vegetation  
578 patches increases biomass production when compared to  
579 simulations where overland flow was not considered. In  
580 contrast, disturbance in the form of overgrazing reduces  
581 positive feedbacks through vegetation and hydrology and  
582 therefore decreases infiltration capacity of potential run-on  
583 cells. The buffer capacity of these hydrological sinks is  
584 reduced, run-off increases and biomass production remains  
585 low. The consideration of overland flow causes water to  
586 flow downstream out of the simulated system and artificial  
587 droughts can occur even in years with good precipitation.

[42] This general process has been observed in other 587  
588 semiarid and arid areas where the spatial exchange between  
589 patches of vegetation affects the resilience of ecosystems  
590 [van de Koppel and Rietkerk, 2004; Ludwig et al., 2005].  
591 Ecosystems heterogeneous in space and linked by spatial  
592 feedback mechanisms provide potential for buffering posi-  
593 tive feedback. Coarse-scale catastrophic shifts are more  
594 likely in systems that have little spatial heterogeneity or

**Table 2.** Comparison of Model Output With Field Data on t2.1  
Percentage Coverage and Biomass Production of Perennial Grasses  
for the Research Area With High and Low Grazing Pressure<sup>a</sup>

	Field Data	Overland Flow	No Overland Flow	t2.2
Cov <sub>L</sub> (%)	50	44 (2.2)	12 (0.9)	t2.3
Cov <sub>H</sub> (%)	1	6 (1.0)	5 (0.6)	t2.4
Prod <sub>L</sub> (t dry matter/ha)	258	262 (5.6)	48 (3.1)	t2.5

<sup>a</sup>Here Cov is coverage and Prod is biomass production, and H is high and L is low grazing pressure. Standard deviations for the model output are given in the parentheses. t2.6



595 no spatial feedback mechanisms to compensate for positive  
 596 feedback. Taking these conclusions into account, our model  
 597 results provide additional insights into the impact of spatial  
 598 exchange of water on ecosystem service and the functioning  
 599 of arid rangelands. Overland flow not only decreases total  
 600 annual productivity but also affects vegetation composition.  
 601 At low grazing intensities, the explicit consideration of  
 602 overland flow favors high abundance of vegetation states  
 603 with high cover of perennial grasses, whereas vegetation  
 604 states with annual vegetation were frequent where overland  
 605 flow was excluded. In contrast, only low effects of overland  
 606 flow on vegetation composition could be observed at high  
 607 grazing intensities (Nabaos). Here, vegetation states consist  
 608 mainly of bare ground, shrubs and, in years of high  
 609 precipitation, annuals, which dominate the simulated land-  
 610 scape at both model versions. Although our results demon-  
 611 strate the high relevance of the explicit inclusion of  
 612 overland flow at landscape scale for vegetation dynamics,  
 613 until now it has seldom been recognized, quantified and  
 614 incorporated into management decisions [Rastetter et al.,  
 615 2003; van de Koppel and Rietkerk, 2004].

616 [43] Generally, rotational grazing strategies have been  
 617 proposed to increase stocking capacity, improve animal  
 618 gains, and improve forage production and range condition  
 619 [e.g., Fynn and O'Connor, 2000]. In this paper, we have  
 620 demonstrated that in semiarid and arid rangelands with  
 621 distinct topography, the rotation of livestock between  
 622 different paddocks combined with moderate stocking rates  
 623 becomes even more important, as nondegraded paddocks  
 624 serve as hydrological sinks which catch and conserve  
 625 surface run-off from degraded paddocks with low vegeta-  
 626 tion cover. A thorough understanding of overland flow is an  
 627 important precondition for improving the management of  
 628 semiarid and arid rangelands with distinct topography.

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